Back-to-Basics Part 1: Developing the CSM & Site Characterization

Understanding Subsurface Fate & Transport

Purpose:

To provide an introduction to the basics of what is necessary to understand contaminant subsurface fate and transport

So you know where contaminant is so you can design a remedy that may actually work
Understanding Subsurface Fate & Transport: Agenda

- Geology
- Hydrogeology
  - Darcy’s law
  - Heterogeneity
  - Diffusion/Dispersion
  - Mass Flux & Mass Discharge
- Contaminant Fate and Transport
  - Contaminant interaction with subsurface, including effects of:
    - pH, DO, and organic carbon content, mineral, and others
    - Role of Anthropogenic activities
- Lessons Learned

Geology

The movement of a substance (both groundwater and contamination) through the subsurface is in great part governed by the physical nature of the subsurface and how the substance interacts with the subsurface.

Therefore, in order to understand contaminant fate and transport, the site manager needs to also understand the site geology.

An important goal of any site investigation should be to at least identify the basic geology of a site and determine its effect on contaminant fate and transport.
Geology

Examples of how the geology affects contaminant fate and transport include

• Coarse grained materials more readily allow transport of contamination
• Finer grained materials may slow contaminant transport and may create contaminant “reservoirs” when contaminants diffuse from coarse grained materials into the fine grained materials.
  • These “reservoirs” are very hard to remediate
  • May act as long term sources of groundwater contamination in flow zones as the contamination migrates from the fine grained materials into the coarser grained materials.
• Very fine grained materials may serve as barriers to contaminant advective migration.

Geology

Examples of how the geology affects contaminant fate and transport include

• Heterogeneities in the subsurface cause contaminants to migrate faster or slower and in different directions than you normally expect, and can create different interactions between the aquifer matrix and contaminants
If you get nothing else from this short course, you must understand that:

Geology controls groundwater flow!

While of course geology controls groundwater flow, past land use can also create significant flow paths and other issues.
Typical landscape in New England has been to a very large extent controlled by the glaciers that covered New England in last ice age:

- Mountains were “smoothed” or sculpted (NH) by glaciers
- Glaciers transported and deposited unconsolidated materials
- Unconsolidated materials have often been reworked by water and anthropogenic activities.

Unconsolidated Geology

Typical New England Unconsolidated Geology

- Till
- Kame terraces/moraines
- End moraines
- Eskers
- Glacial lake and marine deposits
- Reworked by water
Unconsolidated Geology

Till
- Can be very thick
- Very heterogeneous
  - Grain Size
  - Materials
- Can be dry but fractured
- Compacted
- Difficult to drill
- Difficult to remediate

Unconsolidated Geology

Kame terraces/moraines
- Deposited by physical action of the glacier
- Often on the sides of valleys
- Less heterogeneous than till
- Kame terraces may be unsaturated
Unconsolidated Geology

**Eskers**
- Deposited by streams flowing under glaciers
- Can be great aquifers

**Outwash plains**

**End moraines**

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Unconsolidated Geology

**Glacial lacustrine and marine deposits**
- Deposited in water so may be less heterogeneous
- Lacustrine and marine clays
- Deltaic deposits
Unconsolidated materials, especially in valleys, will be reworked by recent alluvial activity (including deposition).

In New England, much of the early industrial development, and many of our cities, are in valleys where materials are often thicker, than in the highlands.
Bedrock Geology

Typical New England Bedrock Geology

- Very Diverse
  - Igneous
  - Metamorphic
  - Sedimentary
- Primary and Secondary porosity
- Structural Considerations
  - Folding/Faulting/Bedding Planes

Intersection of Scale and Fracture Flow Properties

- Macroscopic
- Mesoscopic
- Microscopic
Bedrock Geology

**Macroscopic Flow: The Big Picture**

- Occurs and Regional or Site-wide Scale
- Regional factors influence flow
  - Faults
  - Rivers
  - Changes in lithology
- Remote Sensing and Terrane Analysis to evaluate interaction of multiple structures
  - Orientation, length, connectivity
  - Karst is considered as a whole
  - Overall flow behaving as continuous Darcian flow system
- Knowing how structures interact helps direct investigation at smaller scales

**Mesoscopic Flow: Where we Learn the Most**

- **Plume delineation, flow between multiple wells/boreholes**
  - Orientation, aperture, density, length, and connectivity
  - Influence of matrix characteristics
- **Boreholes and Outcrops**
  - Fracture analysis
  - Hydraulic testing
- **Flow in fracture sets**
  - Impact of turbulent flow may become evident
  - Advection, entrainment, dispersion
- **Primary scale of investigation**
  - Majority of investigation and characterization techniques
Bedrock Geology

Microscopic Flow: Tools for Fine-Tuning your Site Understanding

- Individual fractures to matrix interaction
- Microscopic and individual fracture analysis
  - Individual fracture characteristics
  - Core samples
- Flow between fractures & matrix
  - Changes the morphology of the fracture (Roughness & planarity)
  - Aperture increases or decreases by infilling and dissolution
  - Diffusion and capillary flow
- Interface between fracture and matrix and matrix storage effects

F&T

We may not get down to this scale very often

Site Specific Geology

When you first become involved with a site, initial site specific information on geology may be obtained from:

- Published literature:
  - State geological survey
  - Universities
- Previous site investigation reports
  - For your site
  - For nearby sites
- Aerial Photographs and Maps (geologic and other)
Site Specific Geology

Conduct Site Visit

- Is the site in a valley where unconsolidated materials may be thick, or in the mountains where depth to bedrock is shallow?
- While not specifically geology, what are the present and past land uses?

Develop Initial Conceptual Site Model (CSM)
Conduct Site Investigation and refine CSM

Geology

CSM and Site Investigation: Discussed this afternoon

Hydrogeology Darcy’s Law

Hydraulic Head

\[ h_1 = Z_1 + \frac{P_1}{\gamma} \]
\[ h_2 = Z_2 + \frac{P_2}{\gamma} \]
\[ Z_1 + \frac{P_1}{\gamma} = Z_2 + \frac{P_2}{\gamma} + dh \]

Difference in head, \( dh = h_1 - h_2 \)

\( Z \) = elevation head
\( \gamma = \rho g \) = specific weight of water
\( \frac{P}{\gamma} = \) pressure head
\( \rho \) = fluid density, \( g \) = gravity
Hydrogeology Darcy’s Law

Components of head in a monitoring well:

- **Ground Surface**
- **Potentiometric Surface**
- **Datum (usually seal level)**
- **Point of Measurement (well screen)**

\[ Z = \text{feet/meters above seal level} \]

Datum (usually seal level) \( Z = 0 \)

\[ \frac{P}{Y} \]

\[ h_1 \]

\[ h_2 \]

\[ L \]

Inaccuracies associated with incorrectly installed monitoring wells

- Top of saturated zone
- Screened intervals
Hydrogeology Darcy’s Law

Darcy determined there was a relationship between flow (Q), area (A), and head (h).

The rate of water flow through a tube is proportional to the difference in the height of the water between the two ends of the tube, and inversely proportional to the length of the tube.

From this relationship, it is possible to derive a number of equations that help describe groundwater flow:

\[ v = \frac{Q}{A} = q = -K \left( \frac{h_1 - h_2}{L} \right) = -Ki \]

\[ \frac{Q}{A} = -Ki \quad Q = -KiA \]

\[ q = \text{Darcy flux (L/T)} \quad K = \text{Hydraulic Conductivity (L/T)} \quad i = \left( \frac{h_1 - h_2}{L} \right) \text{ (L/L unitless)} \]
In regards to velocity:
The Darcy equation describes the flow rate per specific unit surface area. It does not consider that flow in the subsurface actually occurs in the effective porosity that is located between the pore grains of the formation. Therefore, the equation has to be changed where $n_e$ is effective porosity:

$$v = \frac{Q}{A} \quad \text{becomes} \quad v = \frac{Q}{A n_e}$$

The flow is also proportional to a coefficient, $K$. In hydrogeology, this needs to be adjusted to account for the formation:

$$Q = K_i A$$

$$k = C d^2 \quad \text{Intrinsic or Darcy permeability (L}^2)$$

$$K = -\frac{C d^2 \rho g}{\mu} \quad K = \text{Hydraulic Conductivity (L/T)}$$

$C = \text{a dimensionless shape}$

$d = \text{average diameter of matrix grains}$

$\rho g = \text{specific weight of fluid}$

$\mu = \text{viscosity of fluid}$
Hydrogeology

Issues with Darcy’s Law

Actual Groundwater flow:

Darcy’s Law Assumes:

1. Saturated and Unsaturated flow
2. Steady-state and transient flow
3. Flow in aquifers and aquitards
4. Flow in homogeneous systems
5. Flow in isotropic and anisotropic media
6. Flow in granular media and fractured rocks

However, Darcy’s law is useful and we can modify it to reflect the real world.

Hydrogeology

High-Energy Deposition

Courtesy Fred Payne, Arcadis
**Hydrogeology**

Where the highest amount of groundwater flow occurs.

- Many soil types are not very conductive
- The conductive soils are laid down in high-energy environments
- High-energy environments are typically heterogeneous and anisotropic

![Soil Conductivity Triangle](image)

Courtesy Fred Payne, Arcadis

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**Representative Elementary Volume (REV)**

- Volumetric dimensions of the scale on which the continuum approach can be used
- Domain of porous media

![Representative Elementary Volume](image)

ITRC ISC-1 Figure 2-3.
Hydrogeology

- Theis (1967): “I consider it certain that we need a new conceptual model, containing the known heterogeneities of natural aquifers, to explain the phenomenon of transport in groundwater.”

Contaminant Fate and Transport

Dissolved Phase Contaminant Transport

Source  Plume  Transects

\[ Q \]

\[ M_d \]

\[ \text{Flux } J_{Bi,j} \]

\[ \text{Flux } J_{A,i} \]

\[ M_{DA} \]

\[ \text{Mass Discharge (} M_d \text{)} = \sum \text{of Mass Flux Estimates} \]

\[ J_{A,i} = \text{Individual mass flux measurement at Transect A} \]

\[ M_{DA} = \text{Mass discharge at Transect A (Total of all } J_{A,i} \text{ estimates)} \]

Photo: USGS

ITRC MASSFLUX-1, 2010
Contaminant Fate and Transport

**Advective Transport:**

In most aquifers, advection is the dominant transport mechanism for dissolved contaminants. The simplest equation used to describe how fast contamination is carried is the equation for the average linear velocity of groundwater:

\[ v_x = \frac{K_i}{n_e} \text{ or } \frac{K}{n_e} \frac{dh}{dl} \]

Where:
- \( v_x \) = the average linear velocity (L/T)
- \( K \) = hydraulic conductivity (L/T)
- \( n_e \) = effective porosity (%)
- \( i \) = gradient \( \frac{dh}{dl} \) (L/L: dimensionless)

**Hydrodynamic Dispersion (Dispersivity)**

However: With an understanding of the role heterogeneity in groundwater flow, diffusion replaces dispersivity as dominant spreading mechanism.
Contaminant Fate and Transport

- Simplifying the subsurface as homogeneous & isotropic has not worked well for remediation-scale plume geometry
- Anisotropy replaces isotropy
- Non-ideal behavior is as pronounced in the vertical

Figure courtesy of Fred Payne, Arcadis

Contaminant Fate and Transport

- As the length scale of interest decreases Diffusion replaces Dispersion in plume behavior
- Geologic heterogeneity and anisotropy also lead to numerous small plumes within each groundwater plume

Figures courtesy of Fred Payne, Arcadis
Molecular Diffusion

Fick’s first law is used to describe molecular diffusion:

\[ J = -D_d \left( \frac{dC}{dx} \right) \]

Where:
- \( J \) = mass flux of solute per unit area per unit time
- \( D_d \) = diffusion coefficient (L²/T)
- \( C \) = solute concentration (M/L³)
- \( \frac{dC}{dx} \) = concentration gradient (M/L³/L) (change in C/change in x)

Fick’s second law is used to describe molecular diffusion when concentrations are changing over time:

\[ \frac{\partial C}{\partial t} = D_d \left( \frac{\partial^2 C}{\partial x^2} \right) \]

Where
- \( \frac{\partial C}{\partial t} \) equals the change in concentration with time (M/L³/T)

I don’t plan on discussing the molecular diffusion equations but if you would like:
Contaminant Fate and Transport

▶ Early time
  • Molecular Diffusion into low permeability zones in the aquifer matrix: "Matrix Diffusion"

▶ Late time
  • "Back Diffusion" out of low permeability zones into higher permeability zones

Mass Flux \( (J) \) is the flux of a contaminant is being carried per unit area of the aquifer.

The one dimensional mass flux equation is:

\[
J = q \times C = \frac{Ki}{n_e} C
\]

Where

\( J = \) contaminant flux per unit area, units of \( \text{M/L}^2/\text{T} \) (mass/area/time)
\( C = \) Contaminant Concentration in units of \( \text{M/L}^3 \) (mass/volume)
\( q = \) specific discharge (\( \text{L}^3/\text{L}^2/\text{T} \))
Contaminant Fate and Transport

Mass Flux Can Be Highly Variable

Flux Sampling Points

Isoconcentration Contours

Transect Wells

Flux Results

Contaminant Concentration

Highest

Lowest

Groundwater Flux

Fast

Slow

Contaminant Fate and Transport

Mass Flux and Concentration

Mass Flux ($J$) = $KiC$

Source Zone

Fine Sand

3%

Mass Flux = 0.03 g/day/m²

K = 1.0 m/day

i = 0.003 m/m

C = 10,000 μg/L

Gravelly Sand

85%

K = 33.3 m/day

i = 0.003 m/m

C = 10,000 μg/L

Mass Flux = 1 g/day/m²

Sand

12%

K = 5.0 m/day

i = 0.003 m/m

C = 10,000 μg/L

Mass Flux = 0.15 g/day/m²

ITRC MASSFLUX-1, 2010
Contaminant Fate and Transport

**Mass Discharge (\(M_d\))** is the integration of the contaminant mass fluxes across a selected transect:

\[ M_d = \int_A J dA \]

A = Area of the control plane in units of \(L^2\)
J – spatially variable contaminant mass fluxes

Can use Mass Discharge of plume to predict constituent of concern concentration in downgradient water supply well

\[ C_{\text{well}} = M_d \div Q_{\text{Well}} \]

\(C_{\text{well}}\) = Concentration in extraction well
\(Q_{\text{well}}\) = Pumping rate for extraction well

Contaminant Fate and Transport

Non-Aqueous Phase Contaminant Transport:

Dense Non Aqueous Phase Liquids (DNAPL)

Light Non Aqueous Phase Liquids (LNAPL)

<table>
<thead>
<tr>
<th>Source Zone</th>
<th>Plume</th>
</tr>
</thead>
<tbody>
<tr>
<td>Phase/Zone</td>
<td>Low Perm.</td>
</tr>
<tr>
<td>Vapor</td>
<td></td>
</tr>
<tr>
<td>NAPL</td>
<td>Vapor Intrusion</td>
</tr>
<tr>
<td>Aqueous</td>
<td>Matrix Diffusion</td>
</tr>
<tr>
<td>Sorbed</td>
<td></td>
</tr>
</tbody>
</table>

*ITRC IDSS-1, Table 2-2 from Sale and Newell 2011*

**KEY POINT:** The 14-Compartment Model helps Stakeholders align on the Life Cycle of the Site and Characterization Objectives
Contaminant Fate and Transport

NAPL Chemical & Physical Properties
- Density
- Solubility
- Viscosity
- Volatility
- Composition

NAPL Interactions with the Sub-Surface Media Affecting Mobility
- Interfacial Tension
- Wettability
- Capillary Pressure
- Saturation
- Residual Saturation

https://www.itrcweb.org/DNAPL-ISC_tools-selection/

Contaminant Fate and Transport

NAPL Saturation

- **Saturation (S)**
  - S is the proportion (percentage) of the pore space occupied by a fluid (NAPL, air, or water)
  - Ranges from 0 to 1.0 (0 to 100%)

- **Residual Saturation (S_r)**
  - S_r is the saturation of NAPL remaining when NAPL is no longer mobile

- **When S < S_r**
  - NAPL will be immobile unless NAPL or solid phase properties change

- **When S > S_r**
  - NAPL may be mobile or
  - NAPL may be potentially mobile but not moving (Pennell et al., 1996, ES&T)

Contaminant Fate and Transport

Relative permeability ($k_r$)

The value of $k_r$ ranges from 0 to 1.0 as a non-linear function of saturation ($S$)

- $k_r$ for groundwater = 1.0 at NAPL $S = 0$
- $k_r$ for DNAPL approaches 1 at as NAPL $S$ approaches 1

(Parker and Lenhard 1987)

Contaminant Fate and Transport

LNAPL Myths

1. LNAPL enters soil pores just as easily as groundwater
2. You can hydraulically recover all of the LNAPL from the subsurface
3. All soil pores in an LNAPL plume are completely filled with LNAPL
4. LNAPL floats on the water table or capillary fringe like a pancake and doesn't penetrate below the water table
5. LNAPL thicknesses in monitor wells are exaggerated (compared to the formation) by factors of 2, 4, 10, etc.
6. LNAPL thicknesses in monitor wells are always equal to the LNAPL thicknesses in the formation
7. If you see LNAPL in a monitor well it is mobile and migrating
8. LNAPL plumes spread due to groundwater flow
9. LNAPL plumes continue to move long after the release is stopped

ITRC 2009  LNAPL 1 Internet Training Module
Contaminant Fate and Transport

**Vertical Equilibrium vs. Pancake Model**
- **Vertical Equilibrium**
  - No LNAPL floats on water table
  - Uniform LNAPL saturation
  - LNAPL penetrates below water table
  - LNAPL and water coexist in pores
- **Pancake Model**
  - Yes LNAPL floats on water table

Attenuation and Retardation of Dissolved Contaminants During Transport

**Focus of this talk is on Organic Contaminants**

**Sorption:**
- Adsorption: reversible, retards contaminant migration
  - chemisorption
  - absorption
  - ion exchange
Attenuation and Retardation of Dissolved Contaminants During Transport

**Sorption:**

\[ C^* = K_d \cdot C \]

Where:

- \( C \) = Concentration of solute in groundwater, mg/l (M/L^3)
- \( C^* \) = Mass of solute sorbed per unit dry weight of aquifer matrix, usually mg/kg (M/M)
- \( K_d \) = adsorption distribution coefficient mL/g or L/kg (L^3/M)

\[ K_d = K_{oc} \cdot f_{oc} \]

Where:

- \( K_{oc} \) = organic carbon partition coefficient, (ml/g [L^3/M]) describes the compound's affinity to organic carbon. \( K_{oc} \) values are available in published literature
- \( f_{oc} \) = fraction organic carbon (%), site/soil specific parameter determined from soil analysis. If it is not possible to analyze soils for \( f_{oc} \), there are tables available provide estimates of general \( f_{oc} \) values for generic soil types.

**Retardation Factor**

\[ R = 1 + \left( \frac{\rho_b K_d}{n} \right) \]

\( \rho_b \) = Bulk Density of the porous media ((kg/m^3 or g/cm^3 [M/L^3]))

Reversible: The amount sorbed to the organic carbon in soil is based on the concentration of the contaminant in water and its affinity for carbon

Maintains plume longevity
Contaminant Fate and Transport

Semi Volatile Organics (SVOC’s)

Generally SVOC’s will rapidly sorb to aquifer matrix.

- Most SVOC’s have high $K_{oc}$ values and will not readily be transported by advective flow
- May however sorb to colloidal particles in groundwater
- When sampling groundwater, in order to understand what mass of contaminants is actually being transported, and what concentrations might be in drinking water, the groundwater sample should not be filtered as filtering will preferentially remove some, but not all of the colloidal particles in groundwater.

Contaminant Fate and Transport

Inorganic Contaminants (Metals)

Generally, metals are relatively immobile in groundwater with typical groundwater chemistry as a result of adsorption, precipitation, chemisorption, or ion exchange reactions onto the aquifer matrix.

- However, changes in pH or REDOX can help mobilize metals
- Metals can also sorb to colloidal partials and be transported by these particles
- Metals can also be found dissolved in liquids discharged to the subsurface and transport with them.
Contaminant Fate and Transport

Inorganic Contaminants (Metals)
- Low flow sampling without filtering is also a must when analyzing groundwater for metals

Inorganic Contaminants (non-metals)
- These include compounds can include Chlorides and Nitrates
- We typically do not address when responding to a release of hazardous materials but they can have some uses.
- As chlorides do not readily sorb to the aquifer matrix, they can be used as tracers
- Nitrates can be an indication of ANFO in groundwater

Contaminant Fate and Transport

At some sites it may be necessary to assess whether a bioremediation remedy is an option:
- pH:
- TOC (total organic carbon)
- REDOX conditions: DO, methane and mineral, and other content
- Bio, sulfate, iron, magnetite
- Oxidant demand
Understanding Fate and Transport

Questions?